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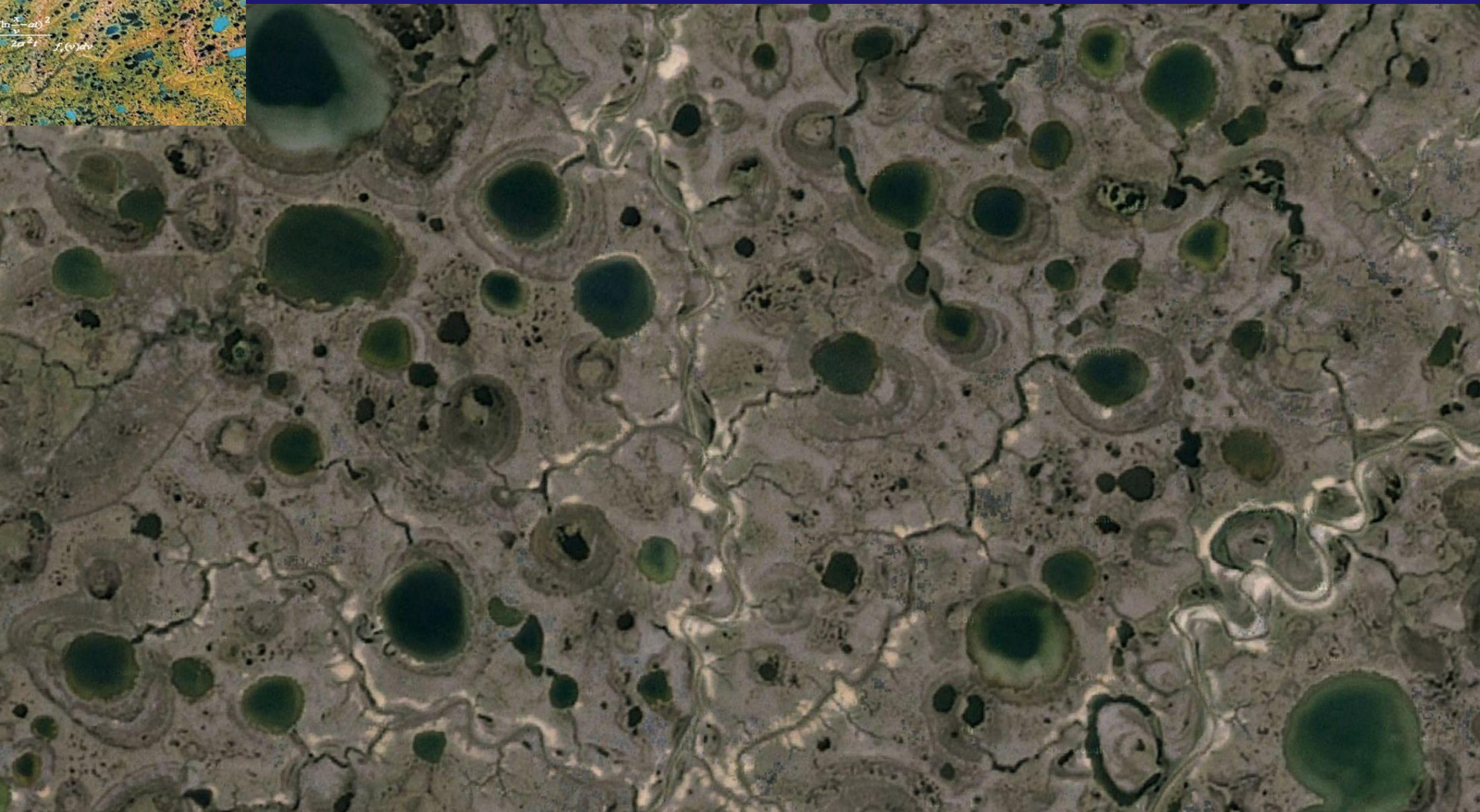


**SERGEEV INSTITUTE OF
ENVIRONMENTAL GEOSCIENCE RAS
(IEG RAS)**

PROBABILISTIC BEHAVIOR MODELING
OF MORPHOMETRIC PARAMETERS
FOR THERMOKARST PLAINS WITH
FLUVIAL EROSION IN
CRYOLITHOZONE

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Lacustrine thermokarst plains with fluvial erosion

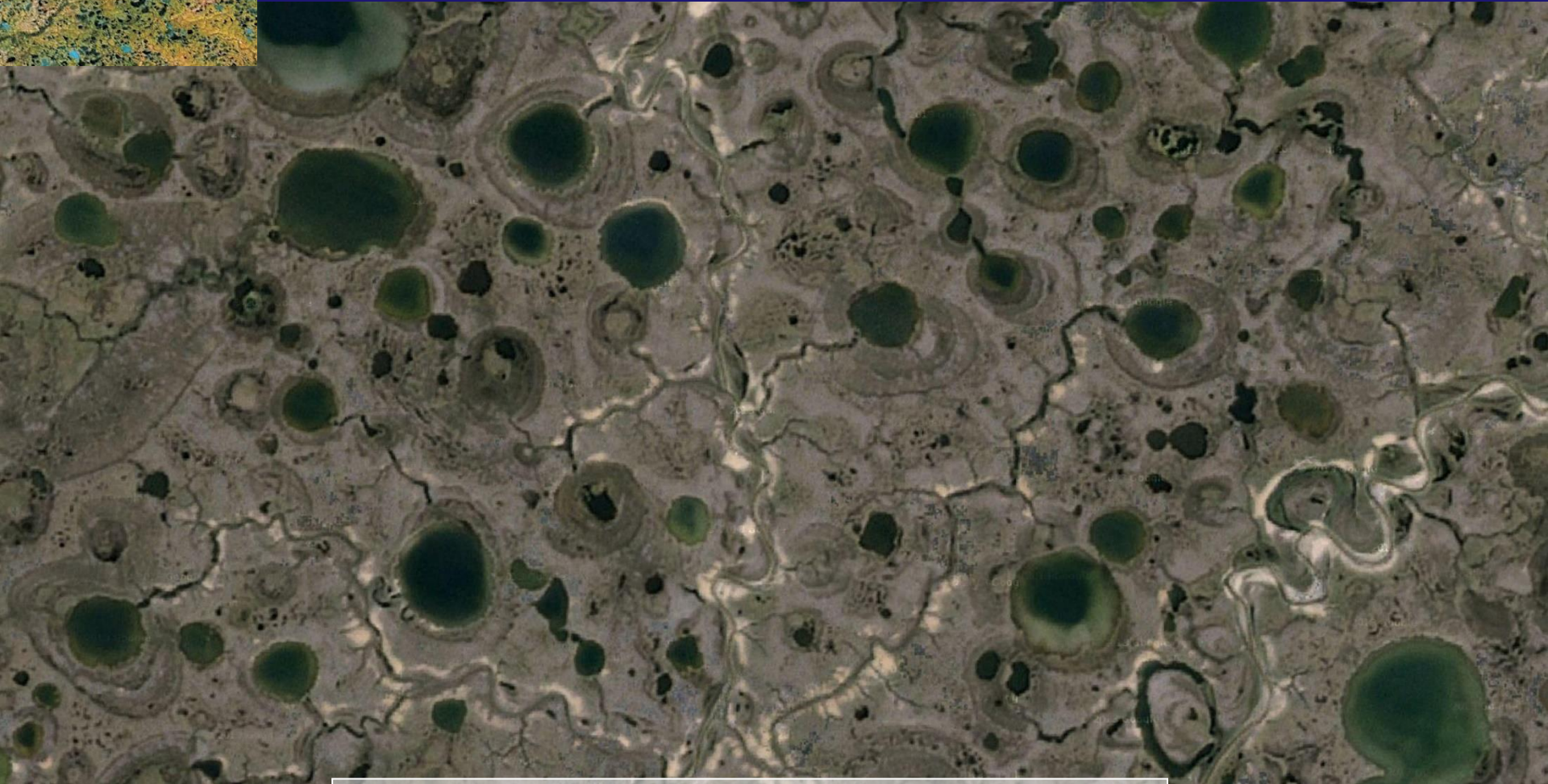


This research aims to consider the patterns of changing morphometric characteristics of thermokarst plains with fluvial erosion and their interrelations.

Thermokarst plains

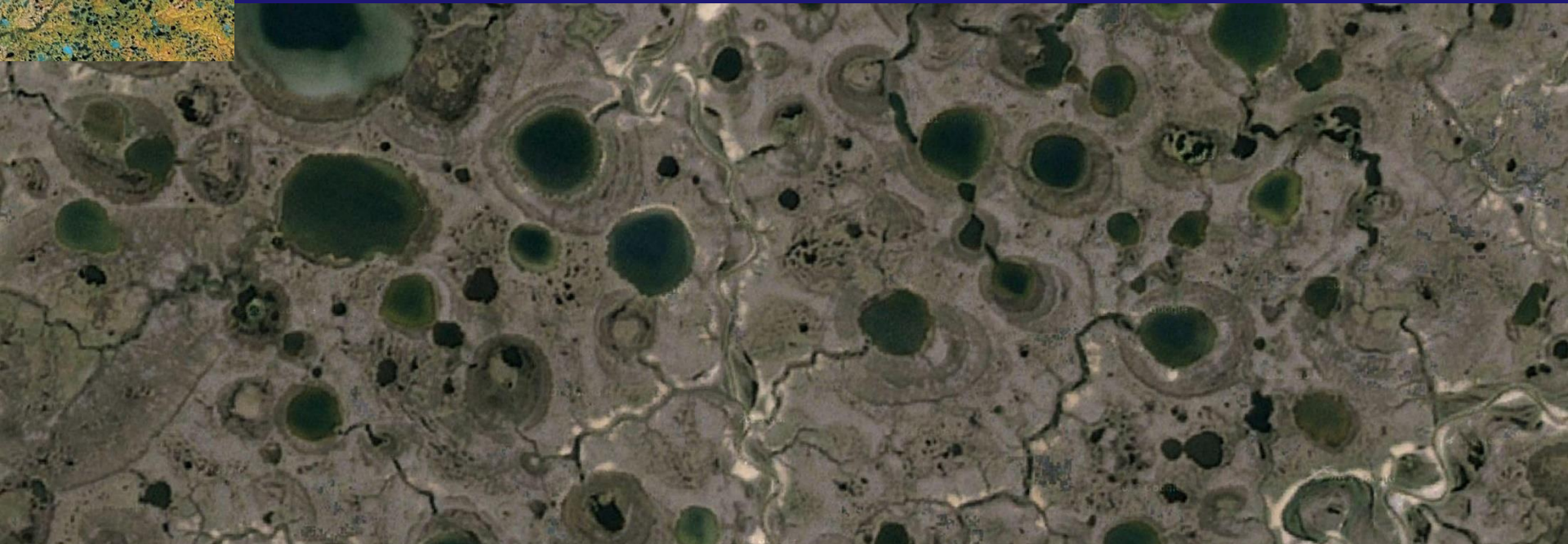


Typical morphometrical parameters of the thermokarst plains with fluvial erosion



**Location density of lakes,
Location density of khasyreis
Average area of lakes
Average area of khasyreis
Lake area percentage
Location density of fluvial sources**

Mathematical model for lacustrine thermokarst plains with fluvial erosion

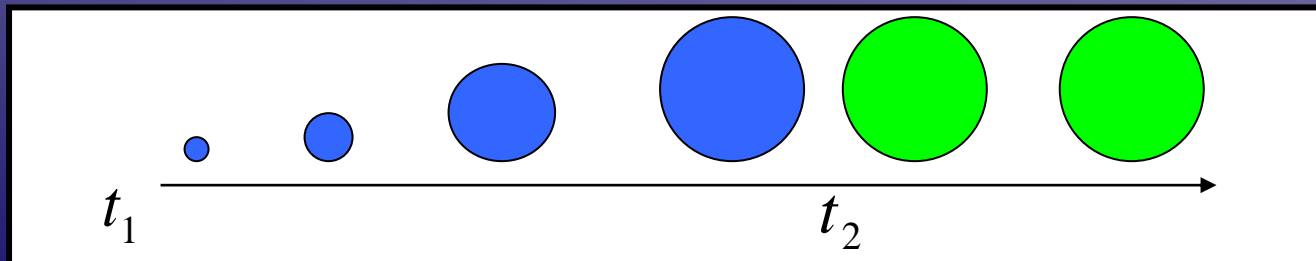
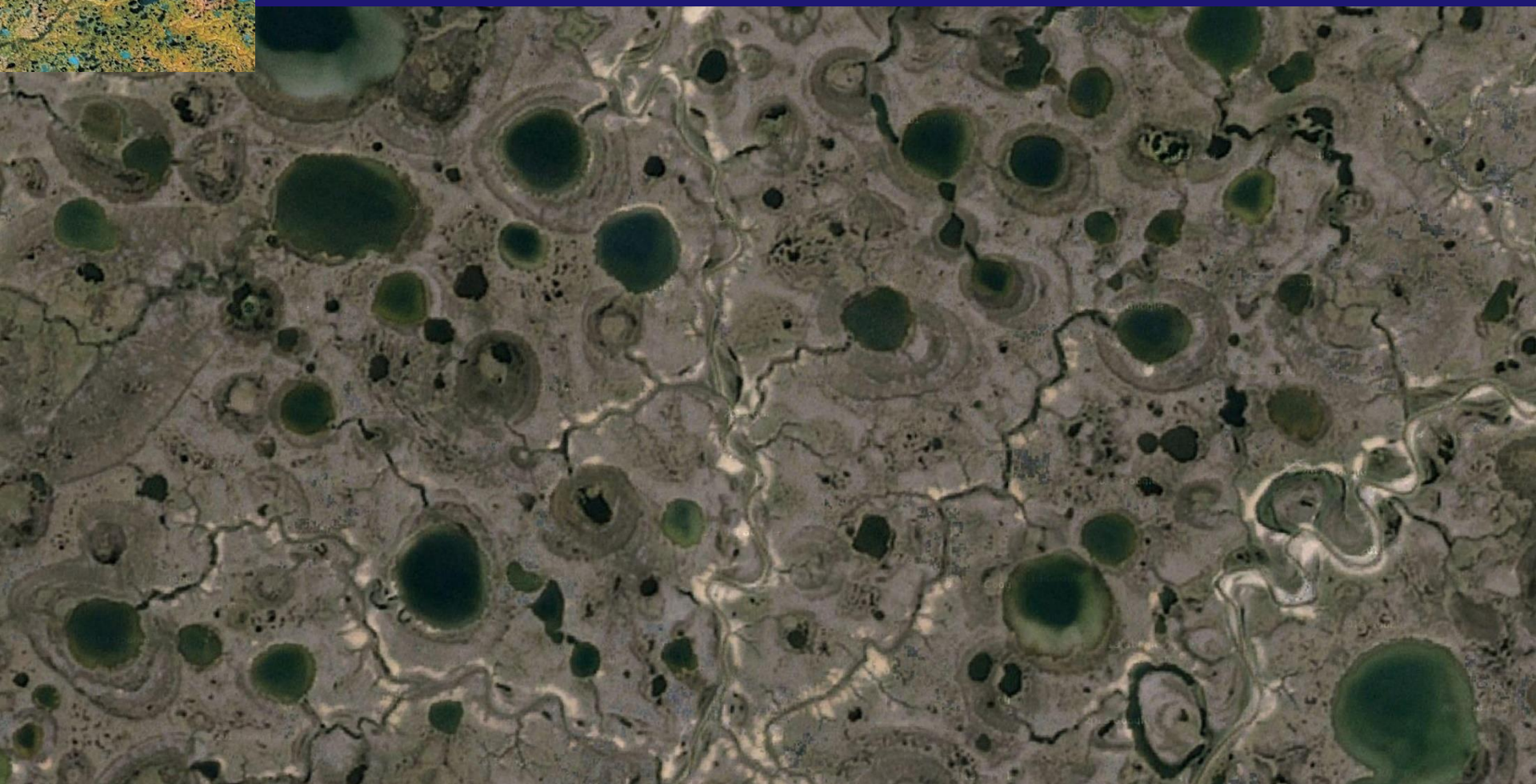


The morphological pattern of these landscapes changes under two opposite processes:

- **appearing and increasing in size of thermokarst lakes,**
- **drainage of the lakes with fluvial erosion and their turn into khasireis.**

Two opposite processes at the same territory is a precondition for a state of dynamic balance.

Mathematical model for lacustrine thermokarst plains with fluvial erosion



Mathematical model for lacustrine thermokarst plains with fluvial erosion.

Model assumptions:

1. Appearance of initial thermokarst depressions (foci) during non-overlapping time intervals (Δt) and non-overlapping sites (Δs) are independent random events. The probability of a depression to appear depends only on the size of the segment and the site

$$p_1 = \lambda \Delta s \Delta t + o(\Delta s \Delta t),$$

where λ is a parameter.

2. Radius of an appeared thermokarst depression is a random variable being a time function; it is undependable of other lakes and the growth rate is directly proportional to heat storage in the lake water.
3. In the course of its growth, a lake can turn into a khasyreï after draining by the erosion network; probability of this does not depend on development of other lakes. If it happens the depression stops to grow.
4. The appearance of new sources of fluvial erosion within a randomly selected area is a random event and its probability depends only on the area.
5. New lakes do not appear within the already existing thermokarst lakes.

Mathematical model for lacustrine thermokarst plains with fluvial erosion. Analysis of the model



$$P(k, t) = \frac{[\lambda t S]^k}{k!} e^{-\lambda t S}$$

The number of initial thermokarst depressions within a trial plot

$$f_0(x, t) = \frac{1}{\sqrt{2\pi\sigma x\sqrt{t}}} e^{-\frac{[\ln x - at]^2}{2\sigma^2 t}}$$

Change of lake size (area, diameter) in case of unlimited growth without taking into account possibility of lake drainage

a, σ - parameters, t - time, S - plot area, λ - generation density of initial thermokarst depressions

Mathematical model for lacustrine thermokarst plains with fluvial erosion. Analysis of the model



$$P(k, t) = \frac{[\eta(t)s]^k}{k!} e^{-\eta(t)s}$$

Distribution of the number of lakes within a random plot at time t

t – time, S – plot area,

$\eta(t)$ - average lake location density at time t

Mathematical model for lacustrine thermokarst plains with fluvial erosion. Analysis of the model



An average number of initial thermokarst depressions appearing per a time unit within a trial plot taking into account total lake area percentage

$$\lambda_1(t) = \lambda[1 - P_l(t)]$$

the length distribution to the nearest fluvial source

$$F(x) = 1 - e^{-\pi x^2}$$

lake radius distribution at moment t

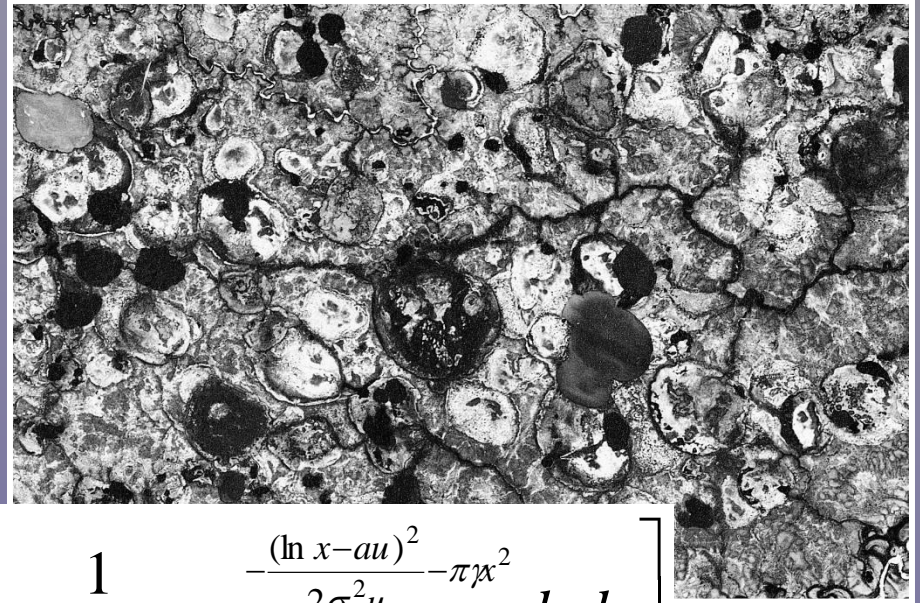
$$f(x,t) = \frac{e^{-\pi x^2} \int_0^t [1 - P_l(u)] f_0(x, t-u) du}{\int_0^t [1 - P_l(u)] \int_0^{+\infty} e^{-\pi x^2} f_0(x, t-u) dx du}$$

a, σ - parameters, t - time, $P_l(t)$ - is a fraction of the whole area occupied by lakes at time t

Mathematical model for lacustrine thermokarst plains with fluvial erosion. Analysis of the model

Dynamics of the lake area percentage

$$P_l(t) = 1 - \exp[-\eta(t) s_l(t)]$$



$$1 - P_l(t) = \exp \left[-\pi\lambda \int_0^t [1 - P_l(u)] \int_0^{+\infty} x^2 \frac{1}{\sqrt{2\pi u x \sigma}} e^{-\frac{(\ln x - au)^2}{2\sigma^2 u} - \pi x^2} dx du \right]$$

where

$\gamma, a, \sigma, \lambda$ - parameters, $P_l(t)$ - is a fraction of the whole area occupied by lakes at time t , $s_l(t)$ - an average lake area, $\eta(t)$ - average lake location density at time t

Mathematical model for lacustrine thermokarst plains with fluvial erosion. Analysis of the model



The main conclusion: after a sufficiently long period of time $t \rightarrow +\infty$ since the start of the process a certain stable state appears under a wide spectrum of conditions: a dynamic equilibrium is established between processes of thermokarst lake origination and their turn into khasyreis.

Limit value of thermokarst affect, which is equal to decision of the equation

$$P_l(t) \rightarrow P_l^* : \ln[1 - P_l^*] = -\lambda\pi[1 - P_l^*]I$$

Limit value for lake location density

$$\eta(\infty) = -\frac{\lambda}{2a} Ei(-\pi\gamma)[1 - P_l^*]$$

Limit value for average lake area

$$s_l(\infty) = -\frac{1}{\gamma Ei(-\gamma\varepsilon)} e^{-\gamma\varepsilon}$$

where a, λ, γ - parameters, $I = \int_{-\infty}^{+\infty} \int_{-\infty}^{+\infty} x^2 e^{-\pi\gamma x^2} f_0(x, u) dx du$

$Ei(x)$ is an integral exponential function, $Ei(-x) = \int_{-\infty}^x \frac{e^{-u}}{u} du$

Mathematical model for lacustrine thermokarst plains with fluvial erosion. Analysis of the model



Limit distribution for lake areas (distribution density) - the integral exponential distribution

$$f_{sl}(x, \infty) = -\frac{1}{xEi(-\gamma\varepsilon)} e^{-\gamma x}, \quad x \geq \varepsilon$$

Limit area distribution for khasyreis – exponential distribution

$$F_h(x) = 1 - e^{-\pi\gamma(x-\varepsilon)}$$

where

γ
 ε

is an average spacing density of erosion sources,

is an initial size of a depression,

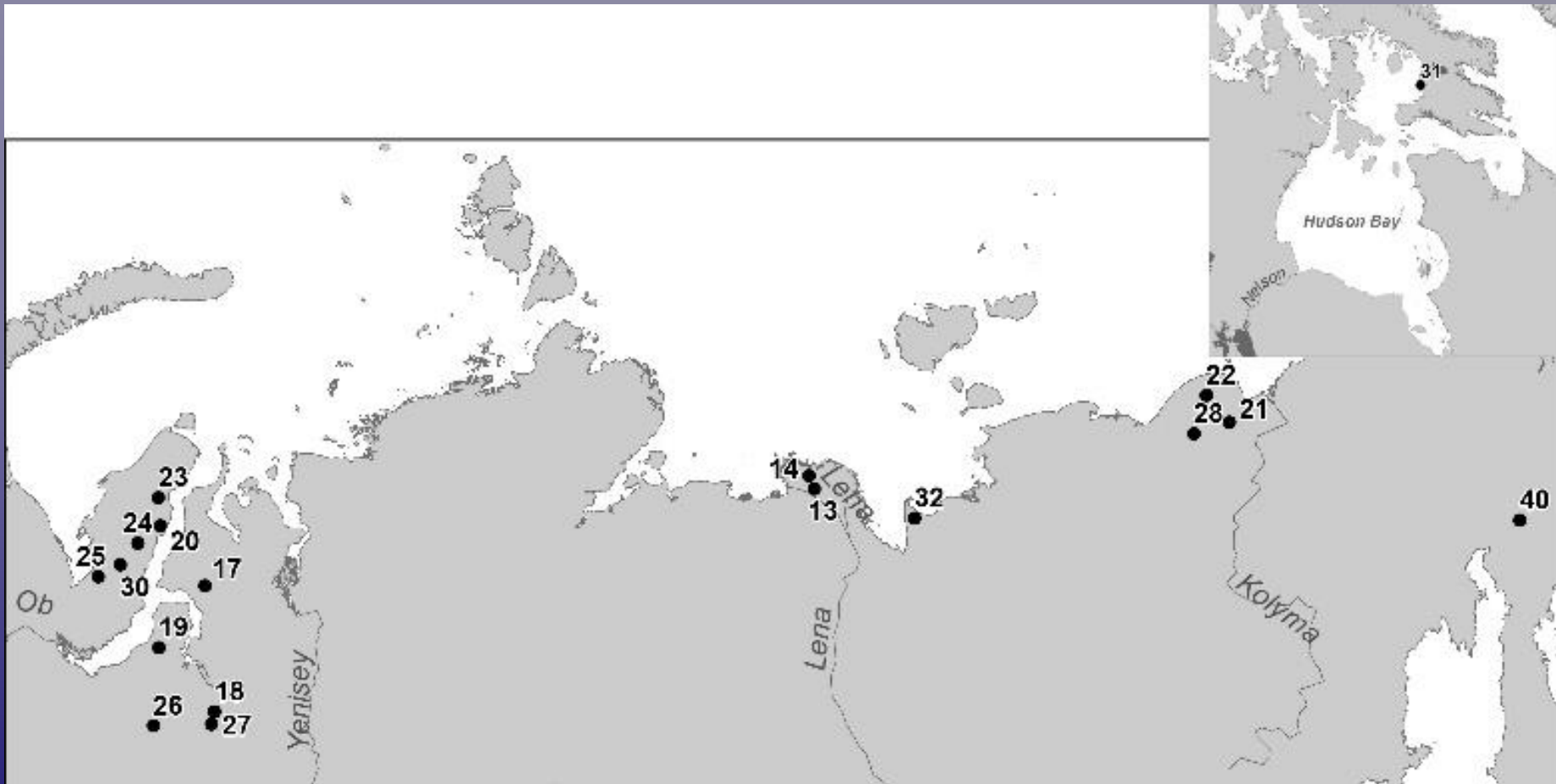
$Ei(x)$

is the integral exponential distribution

$$Ei(-x) = \int_{-\infty}^x \frac{e^{-u}}{u} du$$

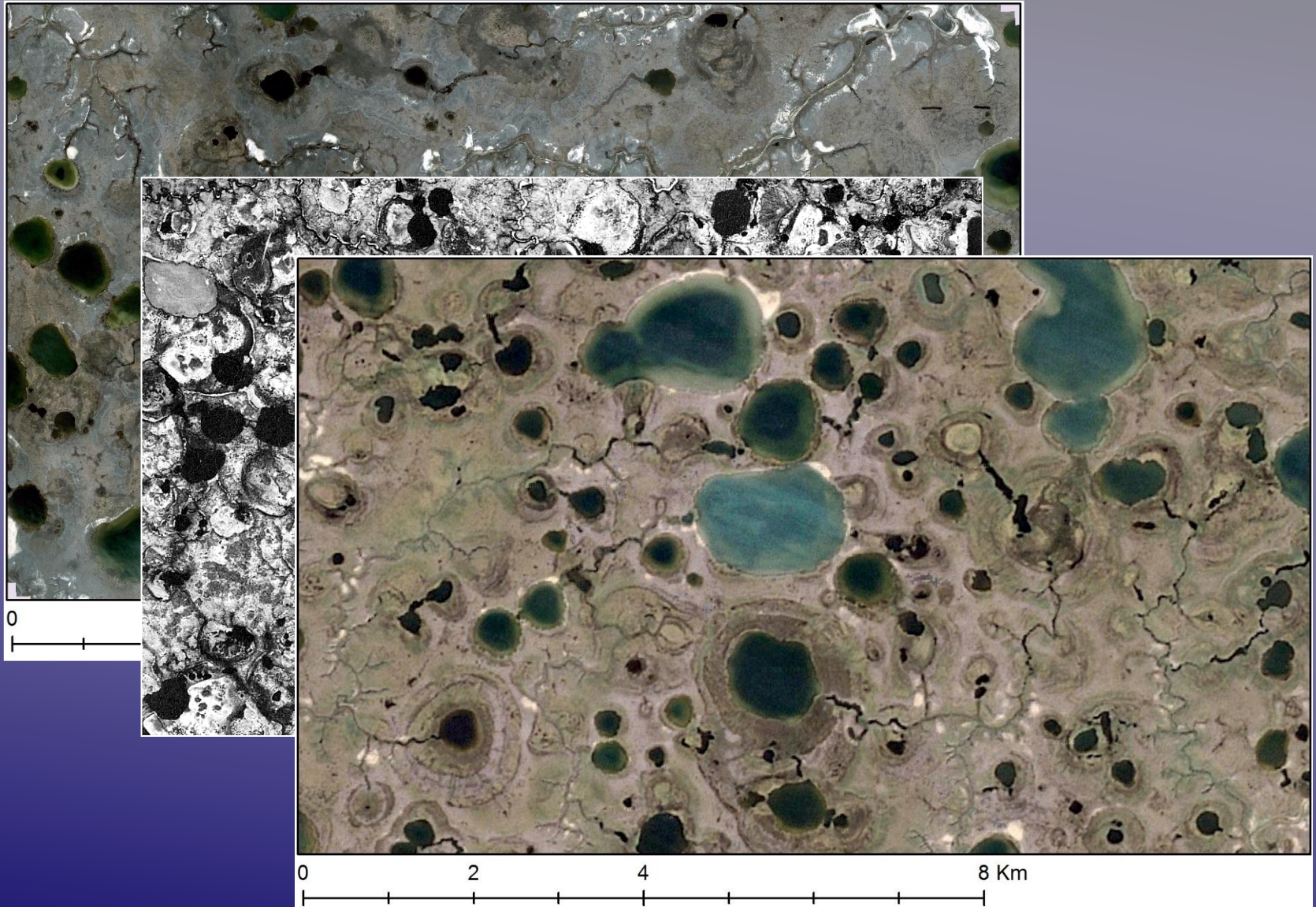
Mathematical for lacustrine thermokarst plains with fluvial erosion. Empirical testing

Key sites located within thermokarst plains with fluvial erosion



Mathematical for lacustrine thermokarst plains with fluvial erosion. Empirical testing

Key sites located in thermokarst plains with fluvial erosion



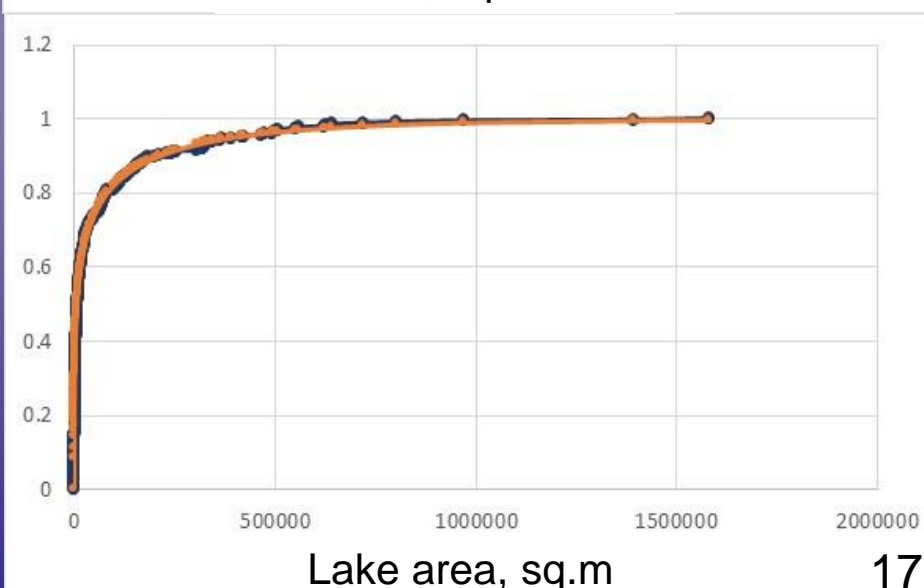
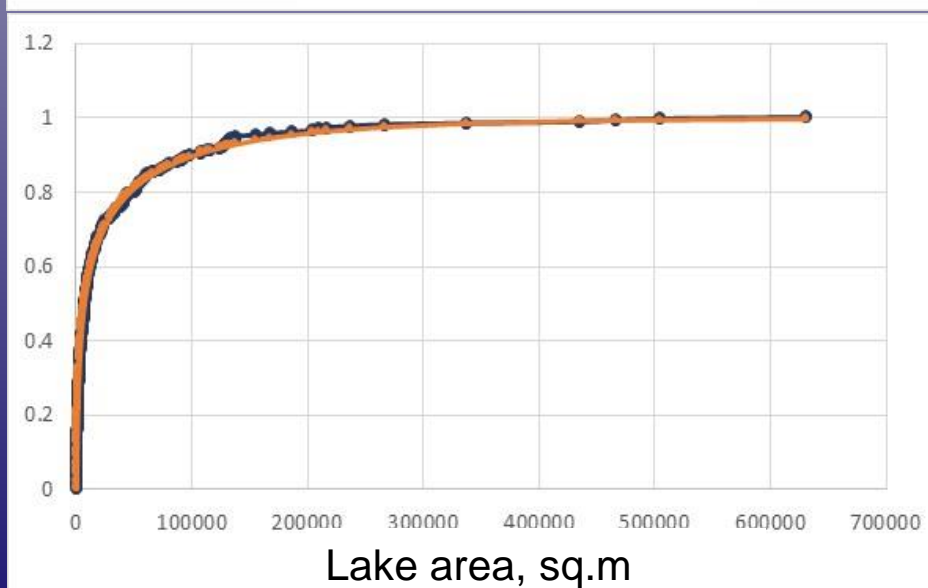
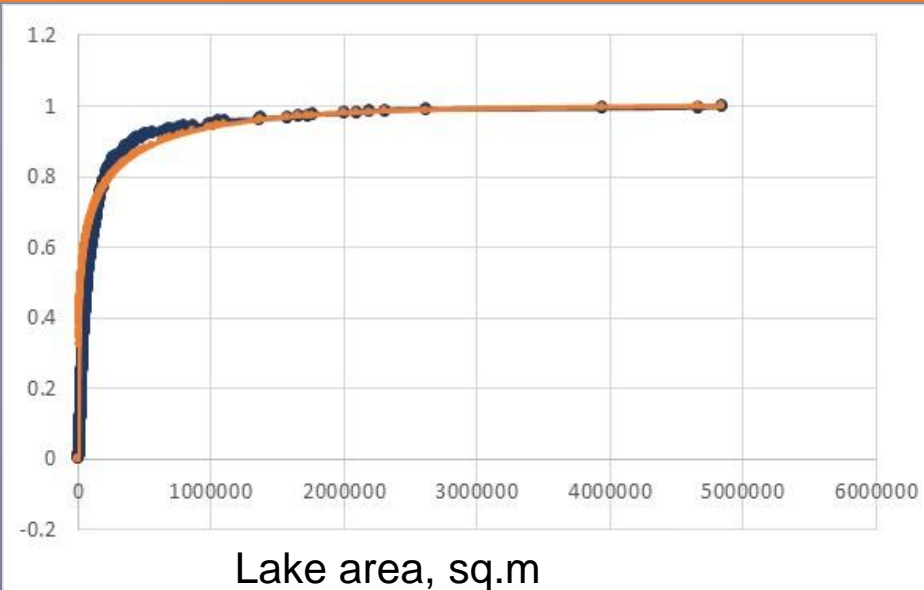
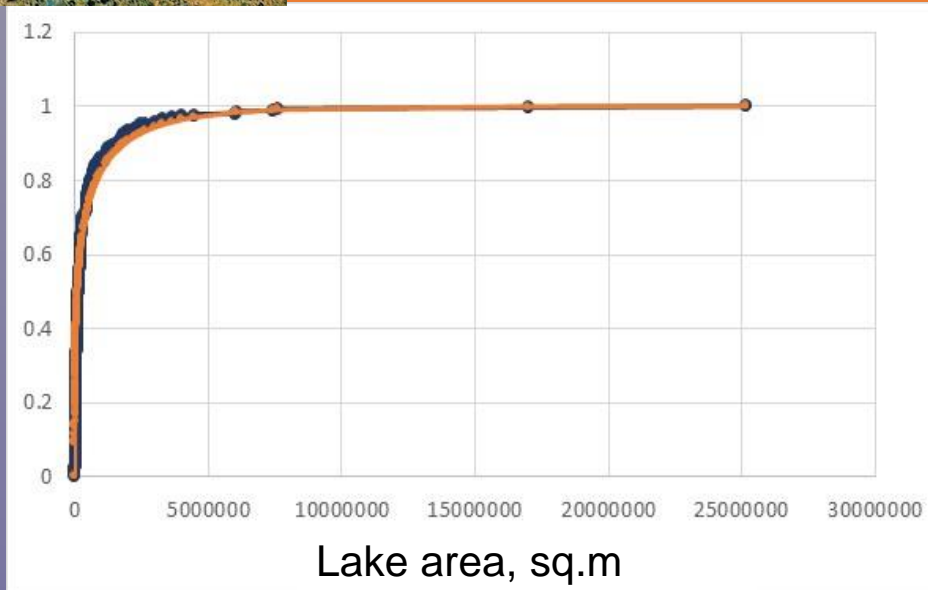
Mathematical for lacustrine thermokarst plains with fluvial erosion. Empirical testing

Correspondence between empirical and theoretical distributions of the average thermokarst lake radii

Key site	Date	13-2	13-1	14-2	17- 2	18- 2	19-2	19- 1	20- 2	20-1	21- 2
sample volume		581	598	209	232	62	161	160	318	359	405
Lognormal		0.000	0.000	0.014	0.005	0.16	0.017	0.091	0.007	0.000	0.000
Gamma		0.000	0.000	0.017	0.000	0.018	0.000	0.000	0.000	0.000	0.000
Normal (for average values)		0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
Integral-exponential		0.000	0.000	0.022	0.002	0.086	0.213	0.394	0.000	0.000	0.109
Key site	Date	13-2	13-1	14-2	17- 2	18- 2	19-2	19- 1	20- 2	20-1	21- 2
sample volume		581	598	209	232	62	161	160	318	359	405
Lognormal		0.010	0.000	0.000	0.044	0.004	0.001	0.225	0.265	0.008	0.310
Gamma		0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
Normal (for average values)		0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
Integral-exponential		0.004	0.014	0.641	0.220	0.663	0.024	0.000	0.000	0.001	0.053
Key site	Date	28-1	30-2	30- 1	31- 2	31- 1	32- 2	32- 1	40- 2		
sample volume		267	519	519	74	70	430	439	535		
Lognormal		0.122	0.322	0.710	0.000	0.000	0.000	0.000	0.001		
Gamma		0.000	0.000	0.000	0.015	0.000	0.000	0.000	0.000		
Normal (for average values)		0.000	0.000	0.000	0.000	0.018	0.000	0.000	0.000		
Integral-exponential		0.085	0.245	0.023	0.005	0.001	0.000	0.000	0.122		

Mathematical for lacustrine thermokarst plains with fluvial erosion. Empirical testing

Correspondence between empirical distributions of the thermokarst lake areas and the integral exponential distribution



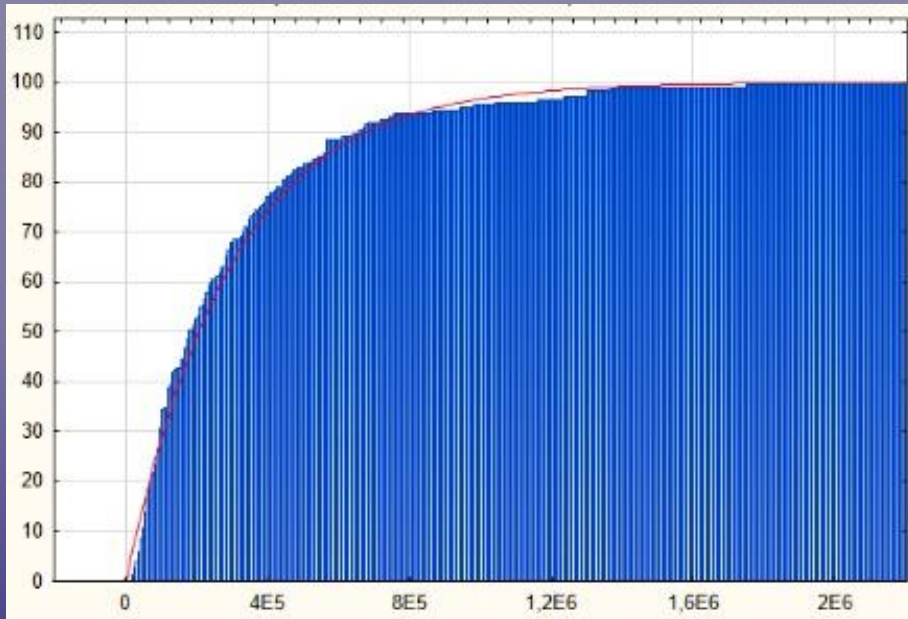
Mathematical for lacustrine thermokarst plains with fluvial erosion. Empirical testing

Correspondence between empirical and theoretical distributions of the thermokarst lake areas

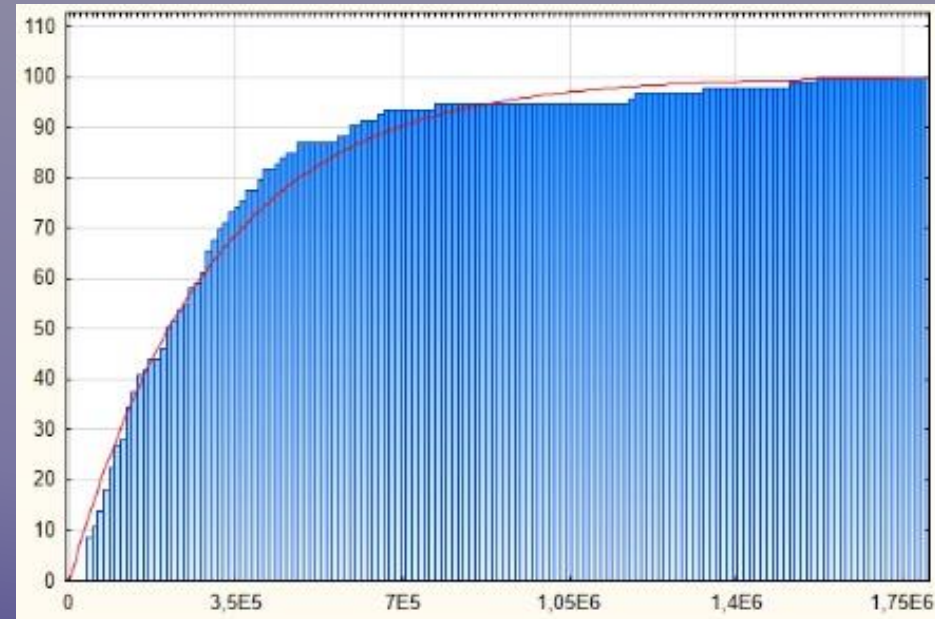
Key site	Date	13-2	13-1	14-2	17- 2	18- 2	19-2	19- 1	20- 2	20-1	21- 2
sample volume		581	598	209	232	62	161	160	318	359	405
Lognormal		0.000	0.000	0.014	0.005	0.16	0.017	0.091	0.007	0.000	0.000
Gamma		0.000	0.000	0.017	0.000	0.018	0.000	0.000	0.000	0.000	0.000
Normal (for average values)		0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
Integral-exponential		0.000	0.000	0.022	0.002	0.086	0.213	0.394	0.000	0.000	0.109
Key site	Date	21-1	22- 2	22- 1	23- 2	24-2	24- 1	25- 2	25-1	26-2	28-2
sample volume		339	244	337	257	346	376	278	281	500	264
Lognormal		0.010	0.000	0.000	0.044	0.004	0.001	0.225	0.265	0.008	0.310
Gamma		0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
Normal (for average radii)		0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
Integral-exponential		0.004	0.014	0.641	0.220	0.663	0.024	0.000	0.000	0.001	0.053
Key site	Date	28-1	30-2	30- 1	31- 2	31- 1	32- 2	32- 1	40- 2		
sample volume		267	519	519	74	70	430	439	535		
Lognormal		0.122	0.322	0.710	0.000	0.000	0.000	0.000	0.001		
Gamma		0.000	0.000	0.000	0.015	0.000	0.000	0.000	0.000		
Normal (for average values)		0.000	0.000	0.000	0.000	0.018	0.000	0.000	0.000		
Integral-exponential		0.085	0.245	0.023	0.005	0.001	0.000	0.000	0.122		

Mathematical for lacustrine thermokarst plains with fluvial erosion. Empirical testing

$$F_h(x) = 1 - e^{-\gamma x}$$



a



b

Graphs demonstrating the correspondence of empirical distributions of the khasyrei areas to the exponential distribution for key sites A3 (a) and A8 (b)



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et al.

Mathematical morphology of
cryolithozone landscape
Moscow
2016

